# GeoStrata

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# Geologic Hazards Assessment Parcel# 21-071-0003 4033 East Nordic Valley View Drive Liberty, Utah

GeoStrata Job No. 1236-002

September 23, 2016

Prepared for:

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#### 1.0 EXECUTIVE SUMMARY

The purpose of this investigation and report was to assess the proposed single family residential building lot located at 4033 East Nordic Valley Drive in Huntsville, Utah for the presence or absence of geologic hazards that might impact the planned development of the site. Hazards assessed in this investigation include landslide, alluvial-fan flooding/debris flow, rockfall, surface fault rupture and stream flooding hazard. Hazards such as slope stability, shallow groundwater, soluble soils, and collapsible or expansive soils were not assessed in this investigation.

The Landslide hazard within the subject site was assessed as part of this study. No landslide related geomorphology were observed and no landslide deposits are reported within or adjacent to the subject lot. The subject site and much of the area south of the subject site is mapped as a landslide and/or landslide undifferentiated from talus, colluvial, rock fall, glacial and soil creep deposits by Elliot and Harty, 2010. However, the subject site is mapped as being underlain by Tertiary Norwood Formation (Tn) by Coogan and King (2016) and Sorensen and Crittenden (1979). It is the opinion of GeoStrata that the landslide and/or landslide undifferentiated from talus, colluvial, rock fall, glacial and soil creep deposits delineated by Elliot and Harty, 2010 is incorrect and is not representative of the surficial deposits within the subject site. Our investigation revealed no indications that the subject lot has been subjected to previous landslides. Therefore, the landslide hazard within the subject site is considered low and it is considered unlikely that landslides will impact the proposed development. It is the opinion of GeoStrata that landslide hazard should not preclude development at the subject lot. It is the opinion of GeoStrata that landslide hazard should not preclude development at the subject lot.

Slope stability of the subject site was not assessed as a part of this geological hazard assessment. Slopes within and immediately adjacent to the subject lot were observed to be gently to steeply dipping approximately 10 to 35 degrees to the north toward Ogden Valley. If it is required that slope stability at the site be assessed, then a site specific geotechnical investigation should be performed for the subject site.

The debris flow and alluvial-fan flooding hazard within the subject site was also assessed as part of this study. No Holocene-aged alluvial fan deposits are mapped within the subject site. Our office and field investigation revealed no indications that the subject lot has been subjected to Holocene-aged debris flows or alluvial fan flooding. The debris flow or alluvial fan flooding

hazards within the subject site is considered low and it is considered unlikely that debris flows or alluvial fan flooding will impact the proposed development. It is the opinion of GeoStrata that debris flow and alluvial-fan flooding hazard should not preclude development at the subject lot.

The rock fall hazard within the subject site was also assessed as part of this study. No rock fall talus or debris resulting from rock fall were observed during our field investigation. Furthermore, sources of rockfall debris were not observed upslope of the subject site. It is our opinion that the rock fall hazard within the subject site is considered low and it is considered unlikely that rock falls will impact the proposed development. It is the opinion of GeoStrata that rock fall hazard should not preclude development at the subject lot.

The surface fault rupture hazard within the subject site was also assessed as part of this study. There are no active faults mapped trending through or within the vicinity of the subject site. The nearest faults to the subject site are the northwest-southeast trending Ogden Valley North Fork Fault and the Ogden Valley Southwestern Margin Fault which are middle and late Quaternary in age (<750,000 years old) with an undetermined reoccurrence interval and a slip rate of less than 0.2 mm/yr (Black and others, 1999). These faults are located approximately 2,020 feet northeast of the subject site and 4,400 feet to the southwest of the subject site, respectively. Therefore, the surface fault rupture hazard within the subject site is considered low and it is considered unlikely that surface fault rupture will impact the proposed development. It is the opinion of GeoStrata that surface fault rupture hazard should not preclude development at the subject lot.

The stream flooding within the subject site was also assessed as part of this study. No streams or drainages were observed within or adjacent to the subject site. Given our field and office investigations, the stream flooding hazard within the subject lot is considered low and it is considered unlikely that stream flooding will impact the proposed development. It is the opinion of GeoStrata that stream flooding hazard should not preclude development at the subject lot. Proper site grading and drainage plans should be developed for the subject site as a part of the civil engineering design for the lot.

NOTICE: The scope of services provided within this report are limited to the assessment of the subsurface conditions for the proposed development. This executive summary is not intended to replace the report of which it is part and should not be used separately from the report. The executive summary is provided solely for purposes of overview. The executive summary omits a number of details, any one of which could be crucial to the proper application of this report.

#### 2.0 INTRODUCTION

# 2.1 PURPOSE AND SCOPE OF WORK

The purpose of this investigation and report was to assess the proposed single family residential building lot located at 4033 East Nordic Valley Drive in Huntsville, Utah for the presence of geologic hazards that may impact the cost and feasibility of the development of the subject site. As part of this assessment, we will identify and describe geologic hazards observed within or immediately adjacent to the subject site. The engineering and design of potential geologic hazards mitigation are out of the scope of this geological hazards assessment. Hazards such as slope stability, shallow groundwater, soluble soils, and collapsible or expansive soils will not be addressed as part of this investigation. If it is required that these hazards be assessed, then a site specific geotechnical investigation should be performed for the subject site.

The work performed for this report was performed in accordance with our proposal and your signed authorization dated September 6, 2016. Our scope of services included the following:

- Review of available references and maps of the area.
- Stereographic aerial photograph interpretation of aerial photographs covering the site area.
- Review of the hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) obtained from the State of Utah AGRC.
- Geologic reconnaissance of the site by an engineering geologist to observe and document pertinent surface features indicative of possible geologic hazards; and
- Evaluation of our observations combined with existing information and preparation of this written report with conclusions and recommendations regarding possible geologic hazards observed to affect the site.

The recommendations contained in this report are subject to the limitations presented in the Limitations section of this report.

### 2.2 PROJECT DESCRIPTION

The subject site is located on the western margin of Ogden Valley just north of Pole Canyon and within the vicinity of the Nordic Valley Ski Area at approximately 4033 East Nordic Valley Way in Liberty, Utah. We understand that the project site is currently an undeveloped single family

residential building lot on a native hillside within the Nordic Valley area. Proposed development, as currently planned, will consist of a single family residential structure as well as associated driveway, utilities and landscape areas. The hillside in the area of the subject lot has moderately steep slopes dipping generally north toward Ogden Valley. It is our understanding that the general area of the subject lot was first developed around the 1960's. The subject site remains in a native condition surrounded by developed residential lots. The subject lot is shown on the Site Vicinity Map included in the Appendix of this report (Plate 1).

#### 3.0 METHODS OF STUDY

# 3.1 OFFICE INVESTIGATION

To prepare for the investigation, GeoStrata reviewed pertinent literature and maps listed in the references section of this report, which provided background information on the local geologic history of the area and the locations of suspected or known geologic hazards (Elliot and Harty, 2010; Coogan and King, 2016; UGS, 2016; Sorensen and Crittenden, 1979). The geologic hazards considered for this site include landslide, alluvial fan flooding/debris flow, rock fall, surface fault rupture and stream flooding. A stereographic aerial photograph interpretation was performed for the subject site using two sets of stereo aerial photographs obtained from the UGS as shown in Table 1.

Table 1

Source	Photo Number	Date	Scale
USDA	AAJ-4FF-8	August 10, 1965	1:20,000
USDA	AAJ-4FF-9	August 10, 1965	1:20,000

GeoStrata also conducted a review of hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) obtained from the State of Utah AGRC to assess the subject site for visible alluvial fan deposits, scarps associated with landslide geomorphology and lineations related to stream flooding hazards or surface fault rupture related geomorphology. The LiDAR elevation data was used to create hillshade imagery that could be reviewed for assessment of geomorphic features related to geologic hazards (Plate 2 Hillshade Map).

# 3.2 FIELD INVESTIGATION

An engineering geologist investigated the geologic conditions within the general site area. A field geologic reconnaissance was conducted to observe existing geologic conditions and to assess existing geomorphology for surficial evidence of geologic hazards. During our fieldwork we conducted site observations to assess geologic hazards that might impact the lot. We used our field observations to confirm the observations made during our office research and to observe any evidence of geologic hazards that were not evident in our office research but which could be observed in the field.

### 4.0 GEOLOGIC CONDITIONS

### 4.1 GEOLOGIC SETTING

The site is located in Huntsville, Utah at an elevation of approximately 5,286 feet above mean sea level along the base of the Wasatch Front Range foothills which border the northwestern margin of Ogden Valley. The Ogden Valley is a northwest trending deep, lacustrine sediment-filled structural basin of Cenozoic age bounded on the northeast and southwest by two normal faults that dip towards the center of the valley. The Ogden Valley is a fault graben flanked by two uplifted blocks, the Wasatch Range on the west and unnamed flat-topped mountains to the east (King and others 2008). The Wasatch Range is the easternmost expression of pronounced Basin and Range extension in north-central Utah (Stokes, 1986).

The near-surface geology of the Ogden Valley is dominated by lake sediments which were deposited within the last 30,000 years during the high stand of the Lake Bonneville Cycle when water inundated Ogden Canyon and formed a small lake in Ogden Valley (Scott and others, 1983; Hintze, 1993; Leggette and Taylor, 1937; King and others, 2008). As the lake receded, streams began to incise large deltas that had formed at the mouths of major canyons along the Wasatch Range and the unnamed flat-topped mountains bounding the eastern margins of Ogden Valley. The eroded material was then deposited in shallow lakes and marshes in the basin and in a series of recessional deltas and alluvial fans. Sediments toward the center of the valley are predominately deep-water deposits of clay, silt and fine sand whereas sediments closer to the mountain fronts are shallow-water deposits of coarse sand and gravel. However, these deep-water deposits are in places covered by a thin post-Bonneville alluvial cover.

# 4.2 SITE GEOLOGY

Surface sediments within the subject site, as shown on Plate 3a Site Vicinity Geologic Map and Plate 4 Site Vicinity 30' X 60' Geologic Map, are mapped as lower Oligocene and upper Eocene Norwood Formation (Tn). The Tertiary Norwood Formation is described as light-gray to light-brown, locally colored light shades of red or green in the west side of Ogden Valley, altered tuff (claystone), altered tuffaceous siltstone and sandstone, and conglomerate (Coogan and King, 2016).

# 5.0 GENERALIZED SITE CONDITIONS

### 5.1 SURFACE CONDITIONS

As stated previously, the project site is located along the base of the Wasatch Front Range foothills and within the western region of Ogden Valley at approximately 4033 East Nordic Valley Way in Liberty, Utah. The subject site is situated on a gently to steeply sloping hillside dipping generally to the north toward Ogden Valley. Slopes in the northern portion of the subject site were observed to be gently dipping approximately 10 to 15 degrees to the north. Slopes in the southern portion of the subject site were observed to be steeply sloping at approximately 20 to 35 degrees to the north. Surficial deposits on the subject site were observed to consist of fine-grained sediment with few up to approximately 4½ feet in diameter well-rounded to subrounded arkosic sandstone and quartzite boulders of the Maple Canyon Formation. The site remains in a relatively natural state, and is thickly vegetated with mature scrub oaks, cedar trees, large native brush and grasses. No structures were observed on the subject property. The properties to the east and west of the subject site are occupied by established residential developments.

## 6.0 CONCLUSIONS AND RECOMMENDATIONS

### 6.1 LANDSLIDE HAZARD

There are several types of landslides that should be considered when evaluating geologic hazards at a site with moderately to steeply sloping terrain. These include shallow debris slides, deep-seated earth or rock slumps and earth flows. Landslides, slumps and other mass movements can develop on moderate to steep slopes where the slope has been altered or disturbed. Movement can occur at the top of a slope that has been loaded by fill placement, at the base of a slope that has been undercut, or where local groundwater rises resulting in increased pore pressures within the slope. Slopes that exhibit prior failures and large landslide deposits are particularly susceptible to instability and reactivation.

Based on review of published geologic maps, our stereographic aerial photograph interpretation, our review of hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) and our field observations, no scarp features, hummocky topography, or other landslide related geomorphology features related to landslide deformation were observed and no landslide deposits are reported within or adjacent to the subject lot (Plate 3a Site Vicinity Geologic Map and Plate 4 Site Vicinity 30' X 60' Geologic Map). The subject site and much of the area south of the subject site is mapped as a landslide and/or landslide undifferentiated from talus, colluvial, rock fall, glacial and soil creep deposits on the Landslide Maps of Utah Ogden 30' X 60' Quadrangle compiled by Elliot and Harty, 2010 (Plate 5 Landslide Hazard Map). However, the subject site is mapped as being underlain by Tertiary Norwood Formation (Tn) as indicated on Plate 3a Site Vicinity Geologic Map and Plate 4 Site Vicinity 30' X 60' Geologic Map. Based on field observations and published geologic maps of the area, it is the opinion of GeoStrata that the landslide and/or landslide undifferentiated from talus, colluvial, rock fall, glacial and soil creep deposits delineated by Elliot and Harty, 2010 is incorrect and is not representative of the surficial deposits within the subject site. Our office and field investigation revealed no indications that the subject lot has been subjected to previous landslides. Therefore, the landslide hazard within the subject site is considered low and it is considered unlikely that landslides will impact the proposed development. It is the opinion of GeoStrata that landslide hazard should not preclude development at the subject lot.

Slope stability of the subject site was not assessed as a part of this geological hazard assessment. Slopes within and immediately adjacent to the subject lot were observed to be gently to steeply dipping approximately 10 to 35 degrees to the north toward Ogden Valley. If it is required that slope stability at the site be assessed, then a site specific geotechnical investigation should be performed for the subject site.

### 6.2 ALLUVIAL FAN FLOODING/DEBRIS FLOW

Alluvial fan flooding is a potential hazard that may exist in areas containing Holocene alluvial fan deposits. This type of flooding typically occurs as a debris flood consisting of a mixture of soil, organic material, and rock debris transported by fast-moving flood water. Debris floods and debris flows can be a hazard on or below alluvial fans or in stream channels above alluvial fans. Precipitation (rainfall and snowmelt) is generally viewed as a debris-flow "trigger", but this represents only one of the many factors that contribute to debris-flow hazard. Vegetation, root depth, soil gradation, antecedent moisture conditions and long term climatic cycles all contribute to the generation of debris and initiation of debris-flows. Events of relatively short duration, such as a fire, can significantly alter a basin's natural resistance to debris-flow mobilization for an extended period of time.

Based on review of published geologic maps, our stereographic aerial photograph interpretation, our review of the hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) and our field observations, no Holocene-aged alluvial fan or debris flow deposit are mapped within or adjacent to the subject site (Plate 3a Site Vicinity Geologic Map and Plate 4 Site Vicinity 30' X 60' Geologic Map). As indicated on Plate 3a and Plate 4, the subject site is underlain by Tertiary Norwood Formation (Tn). Our office and field investigation revealed no indications that the subject lot has been subjected to Holocene-aged debris flows or alluvial fan flooding. Therefore, the debris flow or alluvial fan flooding hazards within the subject site is considered low and it is considered unlikely that debris flows or alluvial fan flooding will impact the proposed development. It is the opinion of GeoStrata that debris flow or alluvial fan flooding hazards should not preclude development at the subject lot.

# 6.3 ROCK FALL

Rock falls are the fastest moving mass movement that predominantly occur in mountains where a rock source exists along steep slopes and cliffs greater than 35 degrees. Rock falls are a result of a loss of support from beneath the rock mass that can be caused by freeze/thaw action, rainfall, weathering and erosion, and/or strong ground shaking resulting from seismic activity. Rockfalls

result in the collection of rock fall material, referred to as talus, at the base of the slope. The presence of talus indicates that a rock fall hazard has occurred and may still be present at the site.

Based on review of published geologic maps, our stereographic aerial photograph interpretation, our review of the hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) and our field observations, no rock fall or talus deposits are located within or immediately adjacent to the subject lot. Boulders and cobbles observed on the subject lot as previously stated, were well-rounded to subrounded and were not observed to be characteristic of a recent rock fall. Sources of rock fall debris were not observed up-slope of the subject site. Our field investigation revealed no indications that the subject lot has been subjected to previous rock fall. Therefore, the rock fall hazard within the subject site is considered low and it is considered unlikely that rock fall will impact the proposed development. It is the opinion of GeoStrata that rock fall hazard should not preclude development at the subject site.

# 6.4 SURFACE FAULT RUPTURE HAZARD

Movement along faults within the crustal rocks beneath the ground surface generates earthquakes. During large magnitude earthquakes (Richter magnitude 6.5 or greater) along the normal faults in the intermountain region, fault ruptures can propagate to the ground surface resulting in a surface fault rupture (Smith and Arabasz, 1991). The fault scarp formed during a surface fault rupture event along a normal fault is generally nearly vertical. A surface rupture fault may be comprised of a larger single surface rupture or several smaller surface ruptures across a fault zone. For all structures designed for human occupancy, a surface rupturing fault is considered active if it has experienced movement in approximately the past 10,000 years (Christenson and others, 2003).

Based on review of published geologic maps, our stereographic aerial photograph interpretation, our review of the hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) and our field observations, no active surface ruptures are located near the subject site (Plate 6 UGS Quaternary Faults Map). The nearest faults are the northwest-southeast trending Ogden Valley North Fork Fault and the Ogden Valley Southwestern Margin Fault which are middle and late Quaternary in age (<750,000 years old) with an undetermined reoccurrence interval and a slip rate of less than 0.2 mm/yr (Black and others, 1999). These faults are located approximately 2,020 feet northeast of the subject site and 4,400 feet to the southwest of the subject site, respectively. Given our field and office investigations, the surface fault rupture hazard within the subject site is considered low and it is considered unlikely that surface fault rupture will impact

the proposed development. It is the opinion of GeoStrata that surface fault rupture hazard should not preclude development at the subject lot.

# 6.5 STREAM FLOODING HAZARD

Stream flooding can be caused by precipitation, snowmelt or a combination of both. Throughout most of Utah floods are most common in spring during the snowmelt. High flows in drainages can last for a few hours to several weeks. Factors that affect the potential for flooding at a site include surface water drainage patterns and hydrology, site grading and drainage design, and seasonal runoff.

Based on review of published geologic maps, our stereographic aerial photograph interpretation, our review of the hillshades derived from the 1 meter Bare Earth LiDAR elevation data (2011) and our field observations, no streams or drainages were observed within or adjacent to the subject site (Plate 7 Drainage Map). Given our field and office investigations, the stream flooding hazard within the subject lot is considered low and it is considered unlikely that stream flooding will impact the proposed development. It is the opinion of GeoStrata that stream flooding hazard should not preclude development at the subject lot. Proper site grading and drainage plans should be developed for the subject site as a part of the civil engineering design for the lot.

#### 7.0 CLOSURE

### 7.1 LIMITATIONS

The conclusions and recommendations contained in this report, which include professional opinions and judgments, are based on the information available to us at the time of our evaluation, the results of our field observations and our understanding of the proposed site development. If any conditions are encountered at this site that are different from those described in this report, our firm should be immediately notified so that we may make any necessary revisions to recommendations contained in this report. In addition, if the scope of the proposed development changes from that described in this report, our firm should also be notified.

All services were completed in accordance with the current standard of care and generally accepted standard of practice at the time and in the place our services were completed. No other warranty, expressed or implied, is made. Development of property in the immediate vicinity of geologic hazards involves a certain level of inherent risk. It is impossible to predict where geologic hazards will occur. New geologic hazards may develop and existing geologic hazards may expand beyond their current limits.

All services were performed for the exclusive use and benefit of the above addressee. No other person is entitled to rely on GeoStrata's services or use the information contained in this letter without the express written consent of GeoStrata. We are not responsible for the technical interpretations by others of the information described or documented in this report. The use of information contained in this report for bidding purposes should be done at the Contractor's option and risk.

### 8.0 REFERENCES CITED

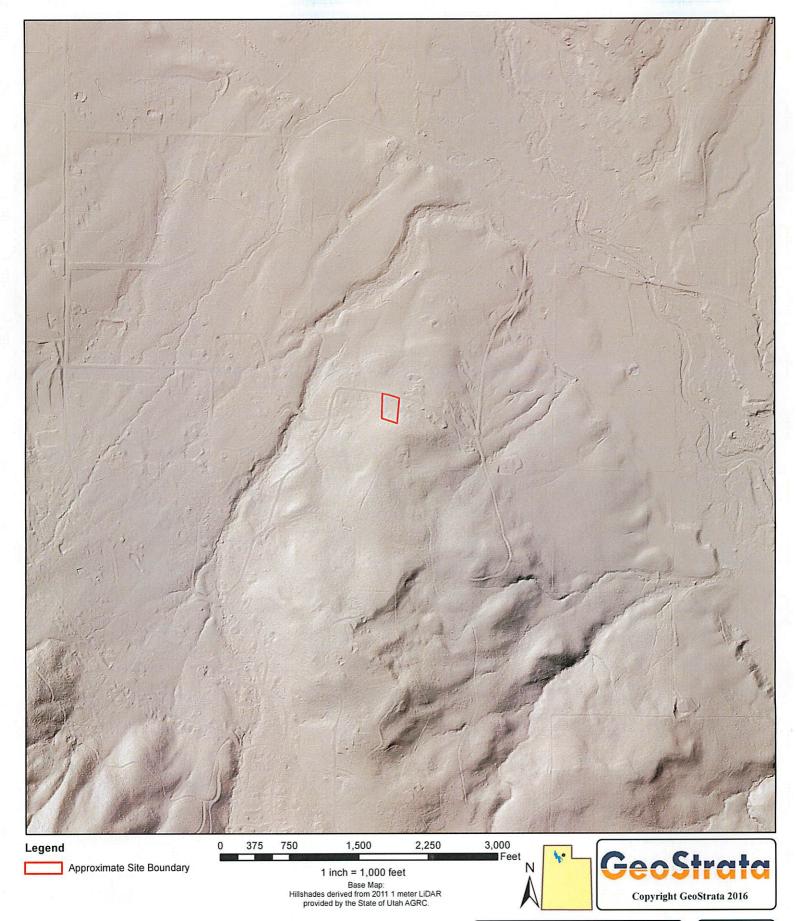
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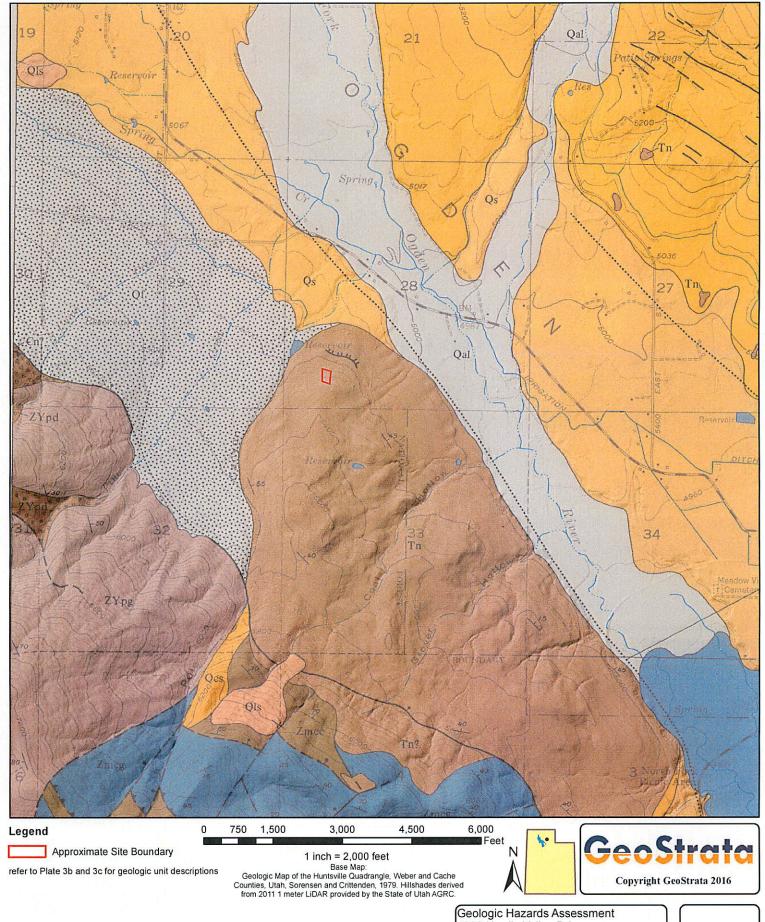
# Appendix



Site Vicinity Map



Hillshade Map



Site Vicinity Geologic Map

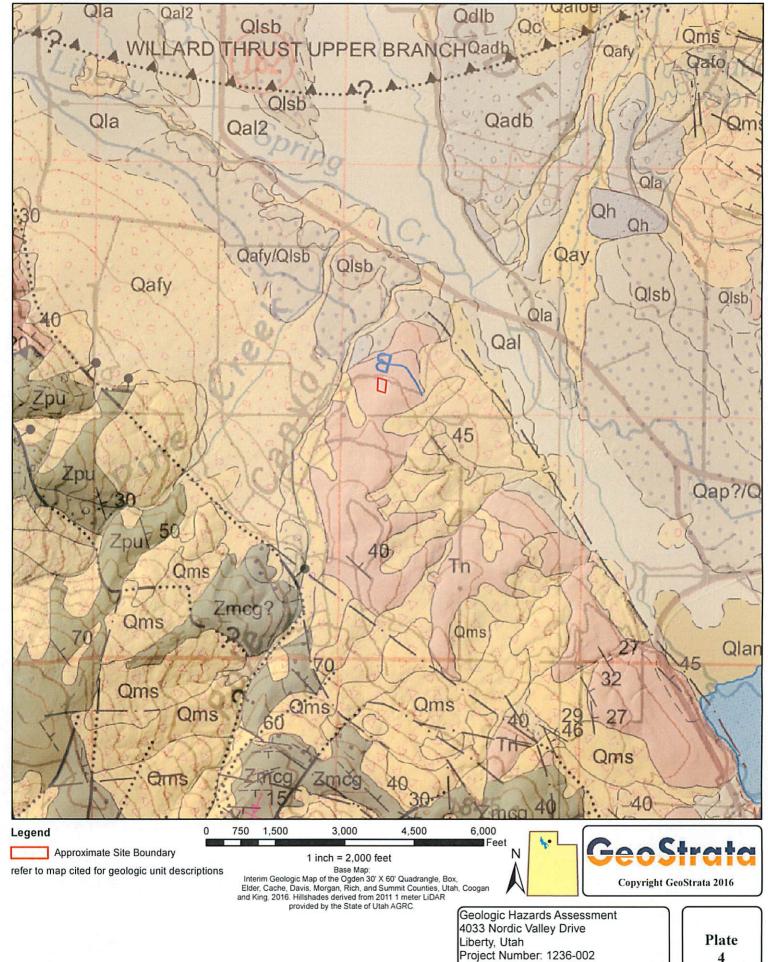
Plate 3a

Qal ALLUVIAL DEPOSITS, UNDIFFERENTIATED (Holocene) remnants of corals and crinoid columnals common throughout unit; Unconsolidated gravel, sand, and silt deposits in presently active thickness 60-69 m stream channels and floodplains; thickness 0-6 m GARDEN CITY FORMATION (Middle and Lower Ordovician) - Thin-COLLUVIUM AND SLOPEWASH (Holocene) - Bouldery colluvium Qcs to medium-bedded, medium- to pale-gray and tan, tan- to buffand slopewash chiefly along eastern margin of Ogden Valley; in part, weathering dolomite, commonly with sandy streaks and lenses. lag from Tertiary units; thickness 0-30 m Interbedded and intercalated with thinly laminated, medium-gray to ALLUVIAL FAN DEPOSITS (Holocene) - Alluvial fan deposits; Qf tan, tan- to buff-weathering siltsone containing nodules and lenses of postdate, at least in part, time of highest stand of former Lake dolomite; thickness 60-75 m Bonneville; thickness 0-30 m ST. CHARLES LIMESTONE (Upper Cambrian) - Includes: Qls €sd LANDSLIDE DEPOSITS (Holocene) - thickness 0-6 m Dolomite member - Thin- to thick-bedded, finely to medium crystalline, light- to medium-gray, white- to light-gray-weathering, TALUS DEPOSITS (Holocene) - thickness 0-6 m cliff-forming dolomite; linguloid brachiopods common in basal 15 m; thickness 150-245 m Qtd TERRACE AND DELTA(?) DEPOSITS (Pleistocene) - In North Fork Worm Creek Quartzite Member - Thin-bedded, fine- to medium-Ogden River, gravel, sand, and silt in stream terraces graded to high grained, medium- to dark-gray, tan- to brown-weathering calcareous stand of former Lake Bonneville; at mouth of Middle and South quartzitic sandstone; detrital grains well-sorted and well-rounded; Fork Ogden River, pinkish-tan sand and silt in delta(?) remnants thickness 6 m deposited during high stands of Lake Bonneville; thickness 0-45 m €n NOUNAN DOLOMITE (Upper and Middle Cambrian) - Thin- to Os SILT DEPOSITS (Pleistocene) - Tan silt and sand forming extensive thick-bedded, finely crystalline, medium-gray, light- to medium-grayflats in Ogden Valley; deposited during high stands of Lake weathering, cliff-forming dolomite; white twiggy structures common Bonneville, but may include older alluvial units; thickness 0-60 m throughout unit; thickness 150-230 m GRAVEL AND COBBLE DEPOSITS (Pleistocene) - In Ogden Canyon, Qg CALLS FORT SHALE MEMBER OF BLOOMINGTON FORMATION Ch gravel and cobble terrace remnants, probably deposited after time of (Middle Cambrian) - Olive-drab to light-brown shale and light- to highest stand of Lake Bonneville; thickness 0-3 m dark-blue-gray limestone with intercalated orange to rusty-brown Qog. OLDER GRAVEL DEPOSITS (Pleistocene) - North of Huntsville, silty limestone; intraformational conglomerate common throughout cobble, gravel, and sand deposit that probably predates high stands unit; thickness 23-90 m of Lake Bonneville; thickness 21 m MAXFIELD(?) LIMESTONE (Middle Cambrian) - Upper part €m NORWOOD TUFF ( lower Oligocene and upper Eocene) - Fine- to thin-bedded, finely crystalline, medium- to dark-gray, ledge-forming Tn medium-bedded, fine-grained, friable, white- to buff-weathering tuff dolomite, often with intercalated light-gray silty limestone; near top and sandy tuff, probably waterlain and in part reworked; thickness of unit, includes distinctive light-gray to white laminated dolomite, 0-450+(?) m underlain by light- and dark-gray mottled limestone. Middle part dominantly olive-drab to greenish-brown micaceous shale, with WASATCH AND EVANSTON(?) FORMATIONS, UNDIVIDED TKwe interbedded medium- to dark-gray limestone, overlain by medium-(Eocene, Paleocene, and Upper Cretaceous?) - Unconsolidated to dark-gray, cliff-forming platy limestone. Lower part dark-bluepale-reddish-brown pebble, cobble, and boulder conglomerate; forms gray, light-gray-weathering, cliff-forming limestone and dolomite, boulder-covered slopes. Clasts are mainly Precambrian quartzite and with intercalated reddish-gray silty limestone; underlain by 30 m are tan, gray, or purple; matrix is mainly poorly consolidated sand thin-bedded, light-blue-gray, slope-forming limestone and shaly and silt; thickness 0-150 m limestone, with some greenish-olive-drab shale. Base of lower unit is LOWER PLATE OF WILLARD THRUST finely crystalline, medium-blue-gray, light-gray-weathering HUMBUG FORMATION (Upper Mississippian) - Medium-bedded, Mb limestone, commonly with intercalated tan to orange-brown silty commonly crossbedded, medium- to fine-grained, gray- to tanlimestone, and locally containing orange-brown oolites near top. weathering quartzite, commonly with thin beds and lenses of Upper and middle parts of formation exposed in Huntsville quadrangle; lower part exposed in North Ogden quadrangle; dark-gray to black chert; interbedded dark- to light-gray mediumbedded dolomite; thickness 300+ m thickness 290 m DESERET LIMESTONE (Upper and Lower Mississippian) - Medium-Md UPPER PLATE OF WILLARD THRUST to thin-bedded, coarsely to finely crystalline, medium-gray- to ST. CHARLES LIMESTONE - See above pale-brown-weathering, dark- to light-gray dolomite and limestone, €sd Dolomite member - See above commonly with thin beds and lenses of dark-gray to black chert; 6 m dark-gray to black mudstone, shale, and oolitic phosphatic shale at Worm Creek Quartzite Member - See above base; thickness 60-75 m GARDISON LIMESTONE (Lower Mississippian) - Upper part finely to Mg £n NOUNAN DOLOMITE - See above coarsely crystalline, thick-bedded to massive, dark-gray- to pale-brown-weathering, medium- to dark-gray fossiliferous dolomite € bo CALLS FORT SHALE MEMBER OF BLOOMINGTON FORMATION with thin beds and lenses of light- to dark-gray chert. Lower part See above finely to medium crystalline, thin- to medium-bedded, commonly €lu CAMBRIAN LIMESTONES, UNDIVIDED (Middle Cambrian) platy weathering, dark-gray to black, light-gray- to blue-gray-Includes limestone and Hodges Shale Members of Bloomington weathering fossiliferous dolomite; thickness 90-260 m Formation, and Blacksmith and Ute Limestones Db BEIRDNEAU SANDSTONE (Upper Devonian) - Medium-bedded to BLACKSMITH LIMESTONE (Middle Cambrian) ) - Medium- to €b laminated, fine- to medium-grained sandstone, dolomitic sandstone, thin-bedded, light-gray to dark-blue-gray limestone; thin-bedded, and dolomite with minor limestone, mudstone, shale, and quartzite; flaggy-weathering, gray to tan silty limestone and interbedded intraformational conglomerate common; weathers to buff, tan, siltstone; light- to dark-gray dolomite, with some reddish siliceous orange, and brown; thickness 75-90 m partings; thickness 400? m HYRUM DOLOMITE (Upper and Middle Devonian) - Thin- to Dh UTE LIMESTONE (Middle Cambrian) - Medium- to thin-bedded, €u thick-bedded, fine- to medium-grained, dark-gray to black, dark- to finely crystalline, light- to dark-gray silty limestone with irregular light-gray-weathering, cliff-forming dolomite; minor intercalated wavy partings, mottled and streaked surfaces, worm tracks, and gray limestone and silty limestone; 5-12 m of medium-grained, bufftwiggy structures common throughout unit; oolites and Girvanella in to tan-weathering dolomitic sandstone locally present in upper 30 m many beds; olive-drab fissile shale interbedded throughout unit. of unit: thickness 107 m Includes thin-bedded, gray-weathering, pale-tan to brown dolomite Dwc WATER CANYON(?) FORMATION (Lower? Devonian) - Thinexposed at base of unit, 18-24 m at head of Geertsen Canyon and bedded to laminated, fine-grained, medium- to pale-gray, pale- to 0-3 m elsewhere; thickness 245? m yellowish-gray-weathering dolomite, silty dolomite, and sandy BRIGHAM GROUP (Crittenden and others, 1971) - Includes: dolomite: thickness 27 m GEERTSEN CANYON QUARTZITE (Lower Cambrian) - Includes: FISH HAVEN DOLOMITE (Upper Ordovician) - Medium- to Ofh Upper member - Pale-buff to white or flesh-pink quartzite, locally thick-bedded, medium to finely crystalline, medium- to light-gray, streaked with pale red or purple. Coarse-grained; small pebbles occur medium- to pale-gray-weathering, cliff-forming dolomite; upper 3 m throughout unit and increase in abundance downward. Base marked weathers very pale gray to silver; small white twiggy structures and by zone 30-60 m thick of cobble conglomerate in beds 30 cm to



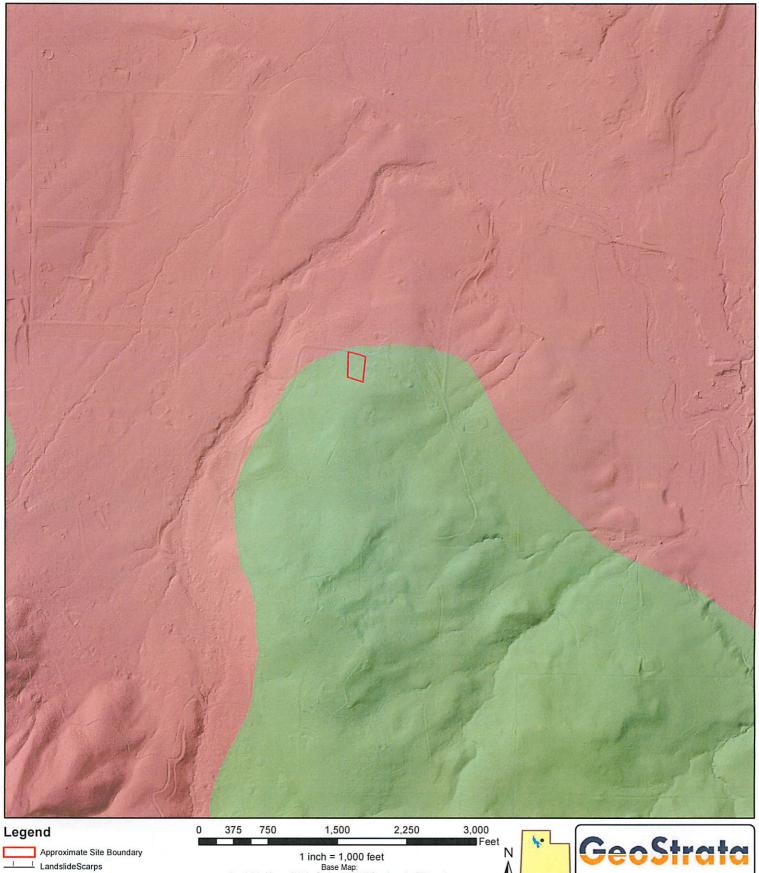
BRIGHAM GROUP (Crittenden and others, 1971) - Includes GEERTSEN CANYON QUARTZITE (Lower Cambrian) - Includes: Upper member - Pale-buff to white or flesh-pink quartzite, locally streaked with pale red or purple, Coarse-grained; small pebbles occur throughout unit and increase in abundance downward. Base marked by zone 30-60 m thick of cobble conglomerate in beds 30 cm to 2 m thick; clasts, 5-10 cm in diameter, are mainly reddish vein quartz or quartzite, sparse gray quartzite, or red jasper; thickness 730-820 m Lower member - Pale-buff to white and tan quartzite with irregular €gel streaks and lenses of cobble conglomerate decreasing in abundance downward, Lower 90-120 m strongly arkosic, streaked greenish or pinkish. Feldspar clasts increase in size to 0.6-1.3 cm in lower part of unit; thickness 490-520 m BROWNS HOLE FORMATION (Precambrian Z) - Includes: Zbq Quartzite member - Medium- to fine-grained, locally friableweathering, well-rounded, well-sorted, terra-cotta-colored quartzite, with some small- to large-scale crossbedding; thickness 30-45 m ~ Zhv ] Volcanic member - Unit comprises volcanic rocks ranging in composition from basalt or andesite to trachyte. Includes gray weathering, fine-grained basaltic flows and a variety of black to red, scoriaceous to amygdaloidal volcanic breecias, all locally reworked as volcanic conglomerate. K/Ar age of hornblende from cobble of alkali trachyte is 570 ± 7 m.y. (Crittenden and Wallace, 1973); thickness MUTUAL FORMATION (Precambrian Z) - Coarse- to medium-grained, commonly gritty, locally pebbly, grayish-red to pale-purple Zm or pink quartzite and feldspathic quartzite with abundant cross bedding; thickness 370 m Zi INKOM FORMATION (Precambrian Z) - Thin-bedded purple and olive-drab to light-green siltstone, argillite and thin-bedded quartzite. Upper half of unit dominantly purple; lower half of unit olive-drab to pale green, and includes thin zone of silver-weathering tuff and sandy tuff; thickness 120 m CADDY CANYON QUARTZITE (Precambrian Z) - Medium-grained, vitreous, white to tan quartzite; unit is dominantly light-colored near top and tan- to pale-brown-weathering in lower part, with abundant intercalated red siltstone at base; thickness 460-600 m Zkc | KELLEY CANYON FORMATION (Precambrian Z) - Upper part interbedded olive-drab siltstone and thin-bedded, tan- or brow weathering quartzite, generally in wavy or contorted beds cut by small sandstone dikelets; contact with overlying unit may be marked by zone of thin-bedded quartzite (0.5-2-cm beds) with redweathering wavy laminae of shale and siltstone. Middle part is gray to lavender argillite enclosing and intercalated with thin-bedded pinkish-gray silty limestone (at Middle Fork Ogden River, shown on map as ls). Lower part is lavender-gray, purple-gray, or olive-drab shale, with thin beds of greenish fine-grained sandstone at top. Base of unit marked by 3-m thin-bedded to laminated, tan-weathering, fine-grained dolomite, thickness 600 m MAPLE CANYON FORMATION (Precambrian Z) - Includes: Zmcc Conglomerate member - Total thickness 30-150 m. Includes: Zmce; Upper conglomerate - Coarse-grained, locally conglomeratic, white quartzite 2mce Argillite - Olive-drab to silvery-gray laminated argillite Zmcc<sub>1</sub> Lower conglomerate - White to pale-gray conglomeratic quartzite, with pebble- to cobble-size clasts of white quartz and white, gray, or pale-pink quartzite Zmcg Green arkose member - Massively bedded pale-green arkosic sandstone, with K-feldspar content locally to 40 percent. Zone of siliceous arkosic quartzite locally present approximately 60 m below top of unit; intercalated quartzitic conglomerates locally present near base of unit; thickness 150-300 m Zmca Argillite member - Olive-drab, locally gray, thin-bedded siltstone and silty argillite, with a medial zone of greenish-gray arkosic sandstone. Argillite commonly shows small-scale folding and marked schistosity. May include rocks of Precambrian Y age near base of unit; thickness 150 m FORMATION OF PERRY CANYON (Precambrian Z or Precambrian Y) ZYpg Gray wacke-siltstone member - Medium- to fine-grained, medium- to dark-gray, tan-weathering graywacke; gray to dark-green, tanweathering, micaceous siltstone; thickness 460 m ZYpd. Diamictite member - Gray to black, tan-weathering diamictite, consisting of pebble- to boulder-size quartzitic and granitic clasts set in a black, medium- to fine-grained sandy matrix; thickness 0-120 m FORMATION OF FACER CREEK (Precambrian X) - Green, purple, Xf and black slate and phyllite; thickness 0-15 m (section incomplete, present only in fault slice) Axial trace of recumbent syncline, showing direction of dip of limbs Approximate location of Lake Bonneville shoreline

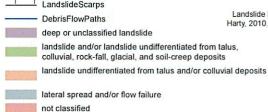




Site Vicinity 30x60 Geologic Map

4





shallow landslide

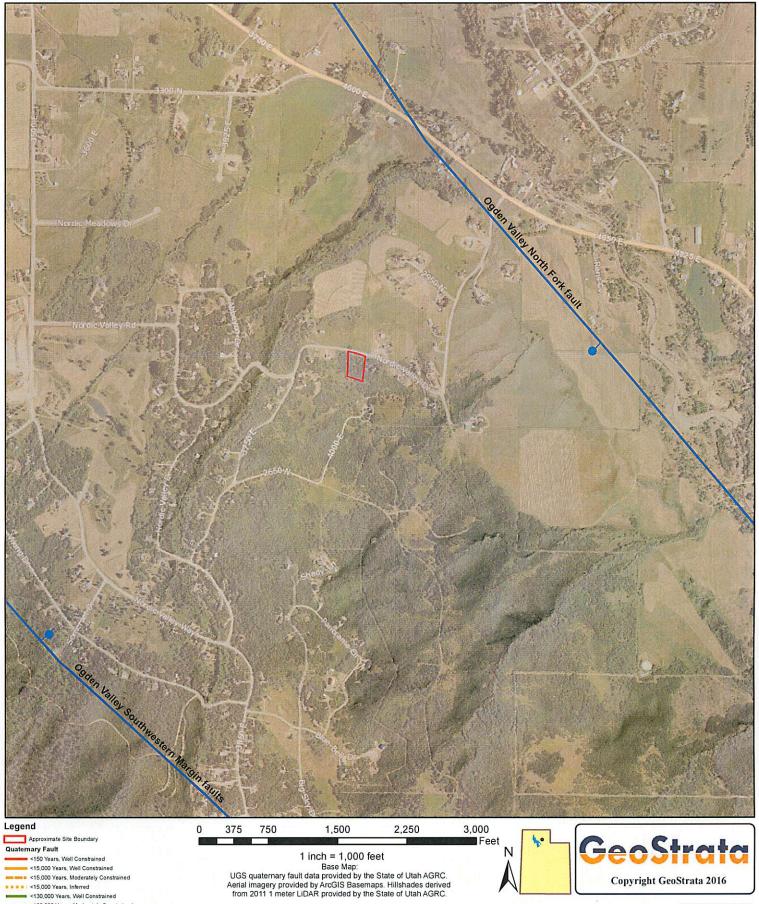
1 inch = 1,000 feet
Base Map:
Landslide Maps of Utah, Ogden 30' X60' Quadrangle, Elliot and
Harty, 2010. Hillshades derived from 2011 1 meter LiDAR provided
by the State of Utah AGRC.



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Landslide Hazard Map



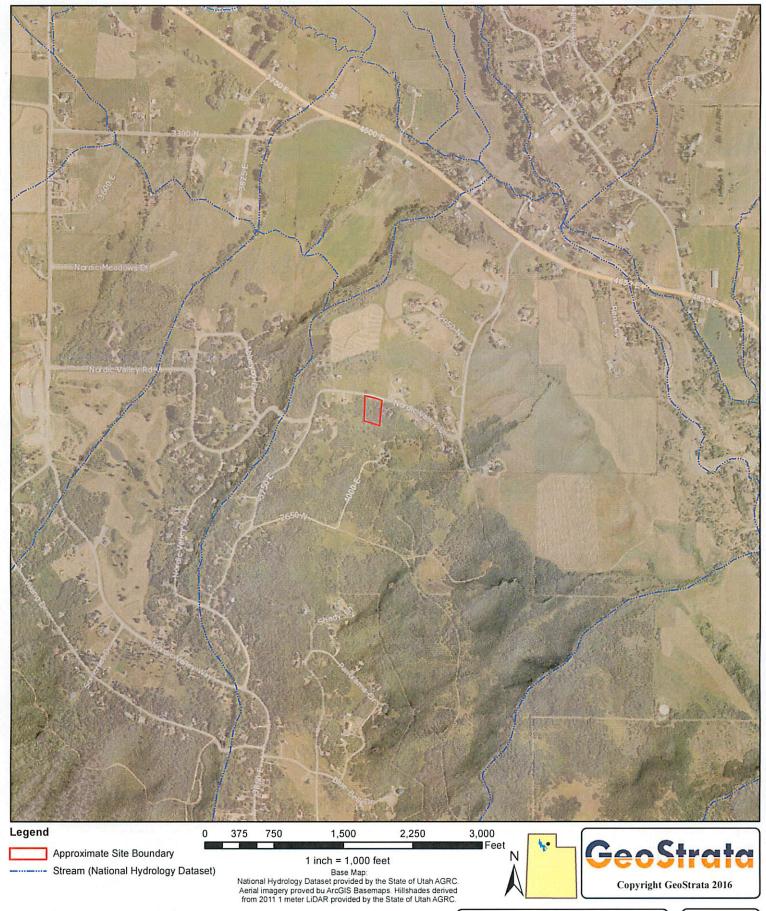
<15,000 Years, Moderately Constrained</p> <15,000 Years, Inferred</p> <130,000 Years, Well Constrained <130,000 Years, Moderately Constrained</p> = = = | <130,000 Years, Inferred <750,000 Years, Well Constrained</p>
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**UGS Quaternary Fault Map** 



Dainage Map